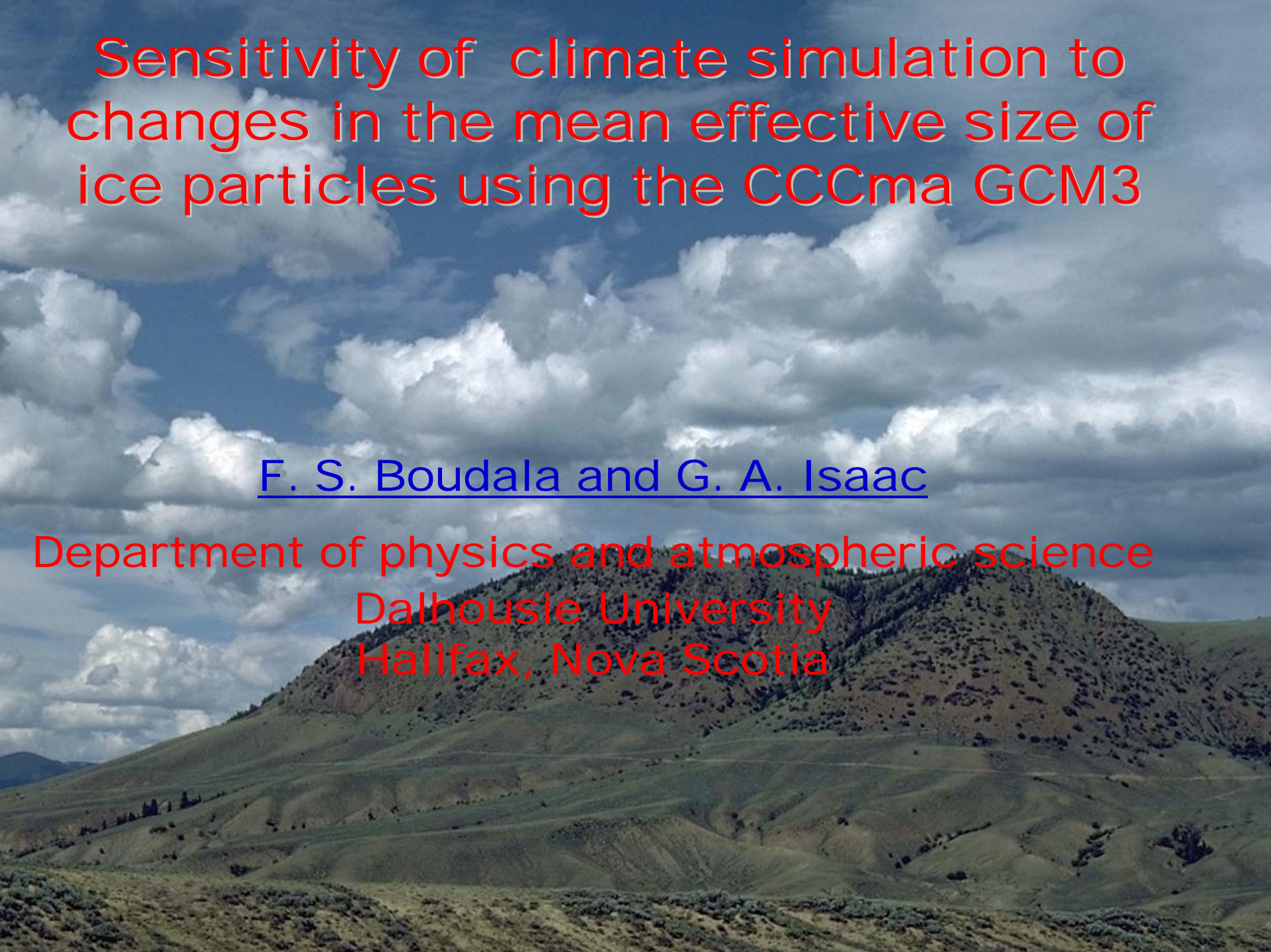


# Sensitivity of climate simulation to changes in the mean effective size of ice particles using the CCCma GCM3

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# ➤ Contents

- ❑ Introduction
- ❑ Parameterization of the mean effective size of ice particles
- ❑ CCCma climate simulation
- ❑ Finally summarize conclusions

# The role of ice clouds

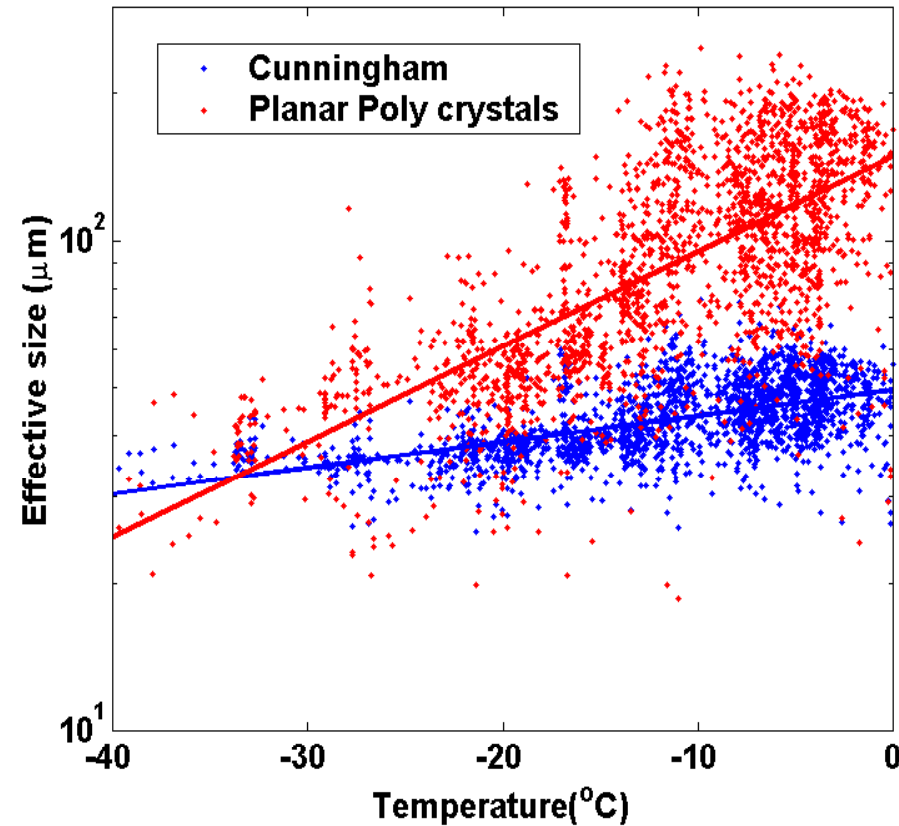
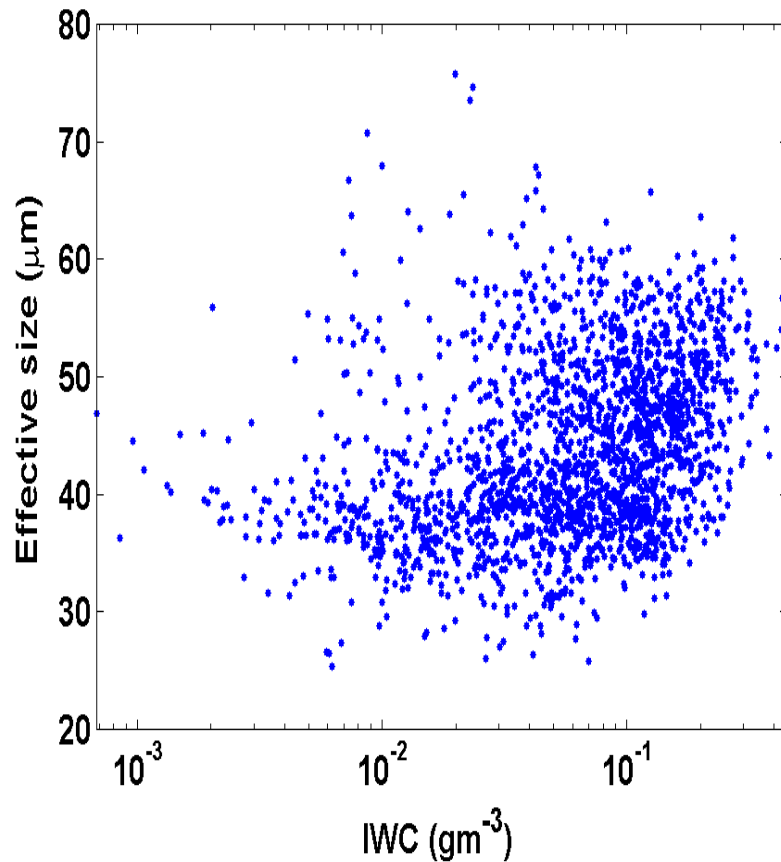
- Effect on climate (no well understood)
  - Due to poor parameterization of the optical properties of ice particles in models
- Climate models are not capable of simulating ice particles SD, thus parameters that are useful for deriving optical parameters such as  $D_{ge}$  and  $A_C$  should be parameterized, but how to parameterize this quantity and how it varies with geographic location is still a contentious issue
- Currently some of the challenging problems are
  - Measurement of small ice particles
  - Measuring or derivation of ice mass in general
- There are some parameterizations of ice particles SD in the mid and Tropical regions, but they suffer all from the same problems.
- In this study Boudala et al. (2002)  $D_{ge}$  parameterizations are tested using the CCCma GCM3

# The definition of the mean effective size of ice particles

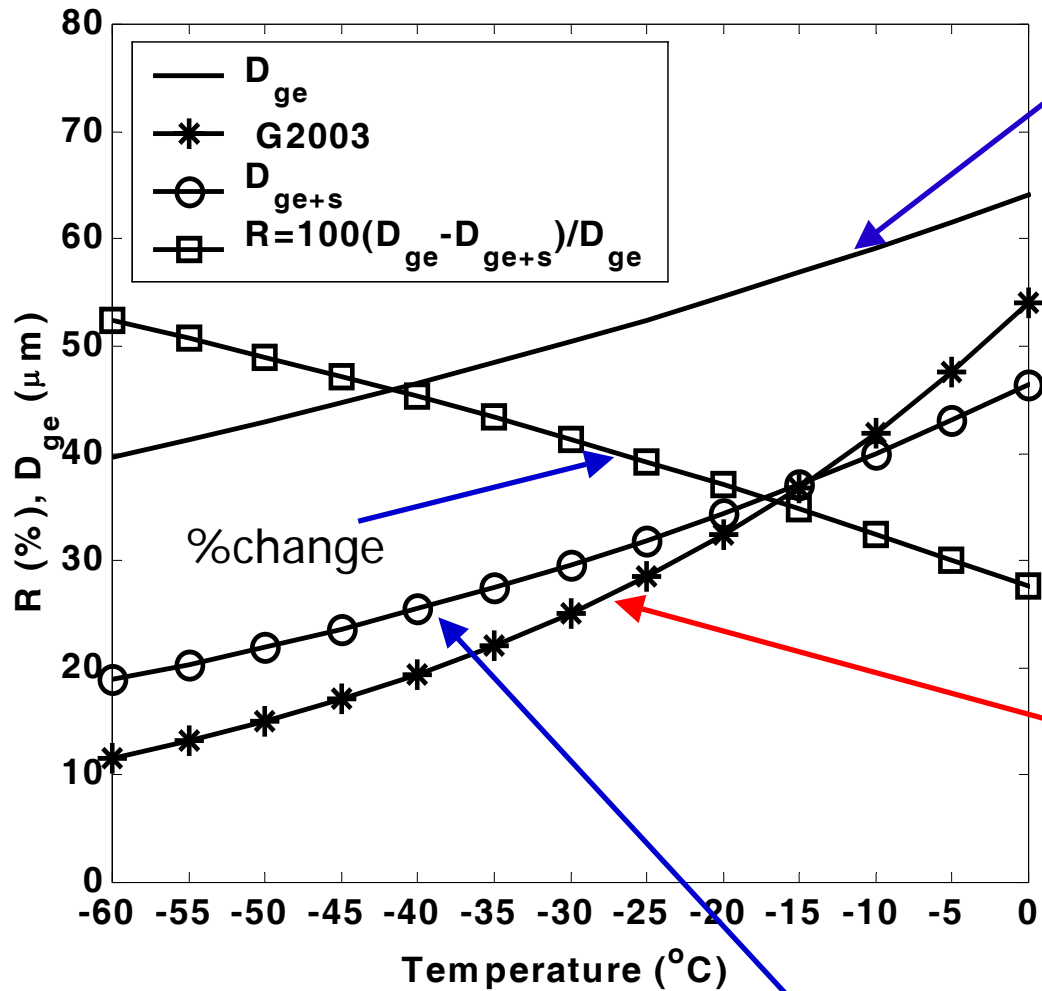
$D_{ge} \sim IWC / \rho_i A_c$  (Fu et al. (1996), others)

- $A_c$  the mean cross-section area per unit volume of ice particles can be measured with optical probes at least for particles greater than 4  $\mu\text{m}$
- The measurement or estimation of IWC and the reliably to determine the density of ice particles are more difficult

# The dependence of the mean effective sizes of ice particles on IWC, temperature, and particle shape



# The temperature dependence of $D_{ge}$ parameterization



Without small ice particles  
Boudala et al. (2002)

Garrett et al. (2003)

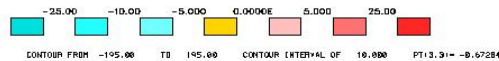
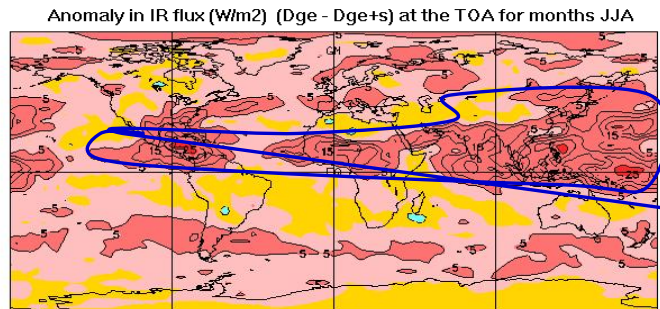
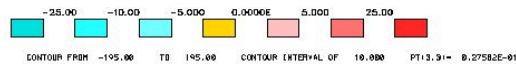
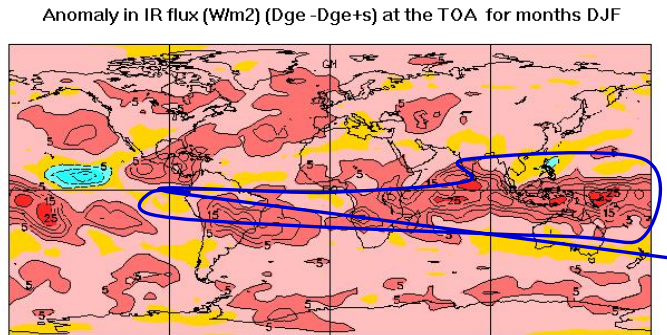
With small ice particles  
Boudala et al. (2002)

# GCM 5 year interactive simulation

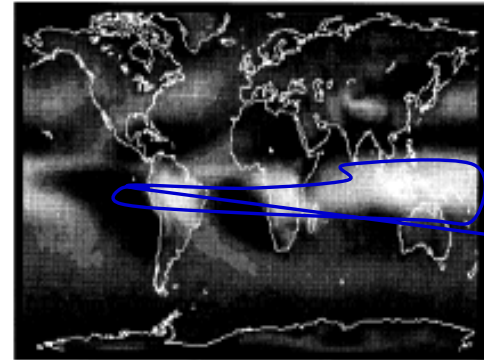
- $D_{ge+s}(T)$  and  $D_{ge}$
- $D_{ge+s}(T, IWC)$
- Optical properties are parameterized as a function of IWC and T
- Cloud water contents are diagnosed based on adiabatic condensation assumption
- Cloud cover is diagnosed based on RH

The OLR anomaly is correlated well with Satellite observation of high clouds (Wylie and Menzel, 1999, *J. Climate*)

## Frequency of high clouds



FREQUENCY OF CLOUDS ABOVE 6 KM  
BOREAL WINTER



BOREAL SUMMER

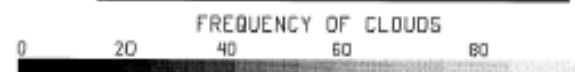
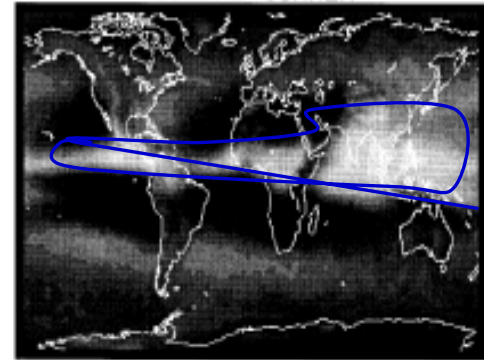
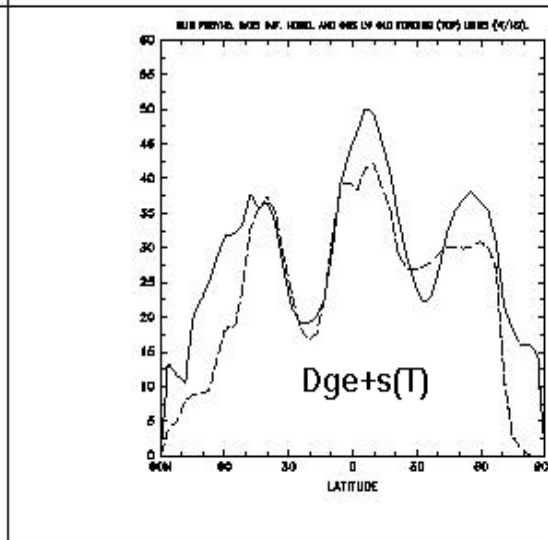
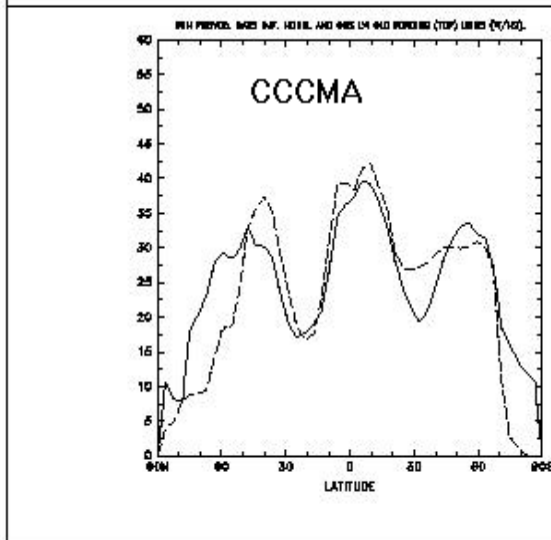
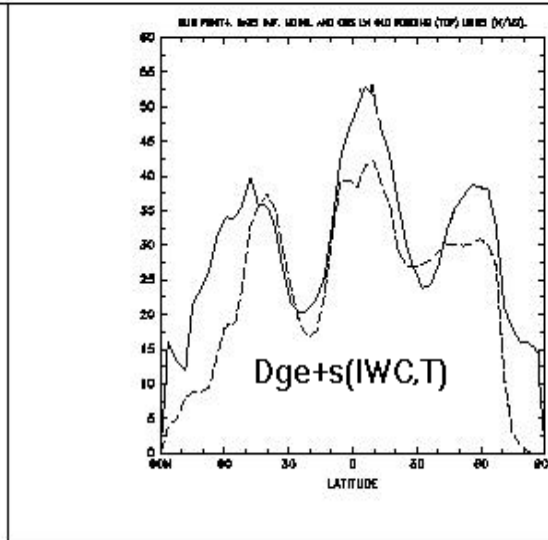
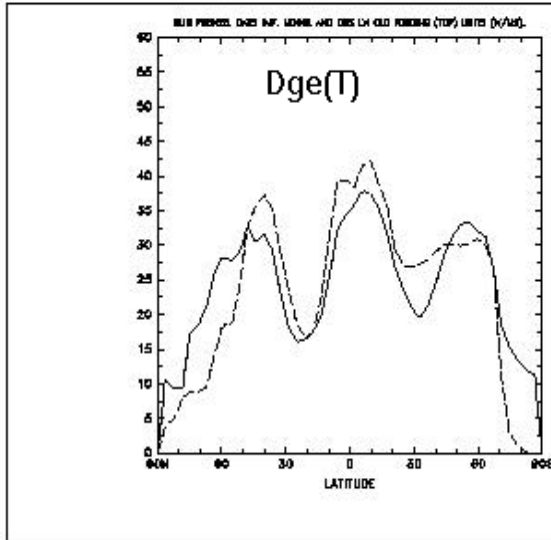


FIG. 3. The frequency of high (above 6 km) cloud detection in a HIRS FUV from 1999 to 1997.

# Zonally averaged long wave cloud forcing at the TOA: Model and satellite in Winter

5 DJFs, long wave cloud forcing



$D_{ge}(T)$  = without small particles

$D_{ge+s}(T)$  = with small particles

$D_{ge+s}(IWC,T)$  = with small particles

The dash line is ERBE satellite observation

- The forcing is more significant in the Tropics with small particles
- For extra-tropical, including small ice particles appear to improve certain regions compared to observations
- The effect of IWC seems to be insignificant

Why is that the OLR anomaly at the TOA is more sensitive in the Tropics?

- The effect of change in clouds?
  - It is possible, but moisture and temperature anomalies are not significant
- One possible explanation could be the change in the  $D_{ge}$  parameterizations with temperature

# Optical effect of small ice particles

$$\beta = \text{IWC}(a_0 + a_1/D_{ge} + a_2/D_{ge}^2), \quad (3.1)$$

$$\beta_s = \frac{\text{IWC}}{D_{ge}}(b_0 + b_1 D_{ge} + b_2 D_{ge}^2 + b_3 D_{ge}^3), \quad (3.2)$$

$$g = c_0 + c_1 D_{ge} + c_2 D_{ge}^2 + c_3 D_{ge}^3, \quad (3.3)$$

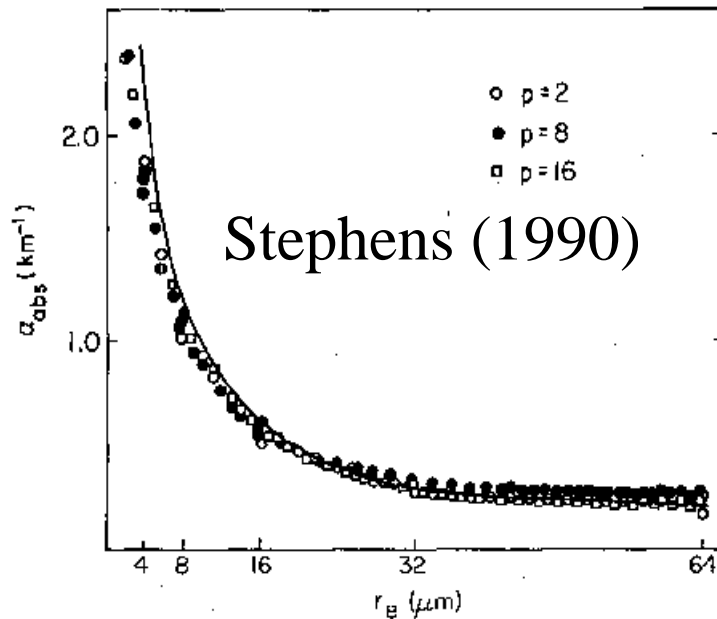
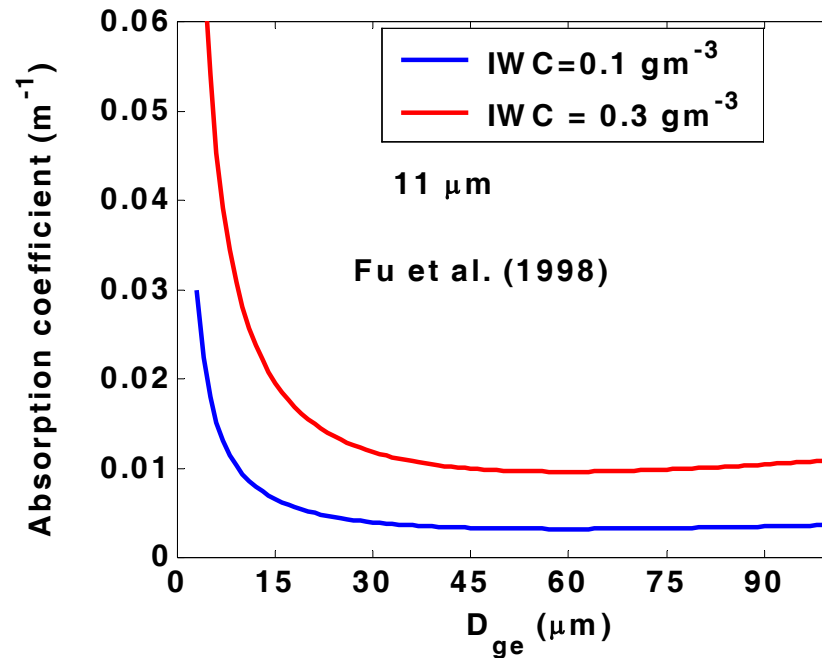


FIG. 3. The relationship between 11  $\mu\text{m}$  volume absorption coefficient and the effective radius as derived from Mie theory using the size distribution given by (17) in the text for the three values of  $p$  indicated and for fixed value of  $w = 0.01 \text{ g m}^{-3}$ .



# The change in global cloud forcing due to small ice particles

## The net cloud forcing $W/m^2$

	with	without	change
Summer	-19	-21.4	2.4
Winter	-20.48	-22.18	1.7

These values are comparable to the current radiative forcing of climate due to an increase in CO<sub>2</sub>, and all other well mixed green house gases since pre industrial time which are estimated to be about 1.46 and 2.43  $W/m^2$  respectively (Ramaswamy et al. 2001).

# Summary of GCM simulation

- The application of  $D_{ge}$  in the current GCMs is uncertain

## Sensitivity to addition of small particles

- **IR radiation**

- Less energy flux at the top of the atmosphere (more absorption)
- More pronounced in the tropics

- **Solar radiation**

- Increased cloud forcing (more reflection)

- **The net effect is spatially variable, but globally induces warming**

- **Model Comparisons of  $D_{ge}(T)$  ,  $D_{ge}(T, IWC)$**

- CCCMA and  $D_{ge}(T)$  gave similar results.
- No significant difference between  $D_{ge+s}(T)$  and  $D_{ge+s}(T, IWC)$ .

- **With observation**

- Hard to compare since CCCMA  $D_{ge}$  is tuned, but generally including small particles seems to improve the model results particularly in the Northern hemisphere mid-latitude, but not in the tropics and the tropics is better captured by  $D_{ge}(T)$  simulation. This suggests that  $D_{ge}$ s in the tropics are larger. But, the recent result by Garrett shows that is not the case. This could be that the CCCma GCM has been tuned with unrealistically large  $D_{ge}$  to match the observations.

***Thank You Very Much***

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# Similar trends in shortwave cloud forcing

## 5 DJFs, Short wave cloud forcing

